DEVELOPMENT OF REFERENCE EVAPOTRANSPIRATION MODELS WITH LIMITED DATA BY USING ARTIFICIAL NEURAL NETWORKS STRUCTURES

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ABSTRACT

The analysis of evapotranspiration (ET) measured at ground level has been mainly considered for agricultural activities, with the aim of determining the water requirements and the supply of water in an appropriate manner. Irrigation activity uses 70% of freshwater in the world, which denotes the importance of this use. ET₀ calculation is a key factor for water balance and crop production, water resource management, irrigation scheduling, and environmental assessment. ET₀ can be obtained with indirect measurements, with high-cost micrometeorological techniques, or can also be estimated with mathematical models. The Food and Agriculture Organization of the United Nations (FAO) has proposed the use of the Penman-Monteith equation as the standard for estimating ET₀ and for evaluating other equations. However, some of the weather variables that are required by this model are often not available in irrigation. In such cases, a simple empirical Hargreaves equation is mainly used. The main objective of this paper is the generation of a model that estimates monthly ET₀ values in Castilla-La Mancha region (Spain) from available termo-pluviometric weather stations with monthly data that can be applied to historical time series. The procedures to estimate the ET₀ values in this study are in one hand, the use of Hargreaves equation, which requires the maximum and the minimum monthly temperature differences, together with extraterrestrial global solar radiation data; and on the other hand, the application of Artificial Neural Network (ANN) structures with the same climate data. Calibration and validation of the ANN models are performed by using data from the years 2000 to 2008 of 44 complete agroclimatic weather stations. The monthly average of the Penman-Monteith daily ET₀ is compared with the monthly values of ET₀ using Hargreaves and the ANN models. The obtained relative errors are: 1) spatial error (spatial validation) of 17.0% and 13.21% for Hargreaves equation, and ANN model respectively, 2) temporal errors (temporal validation) of 13.8% and 12.0% for Hargreaves equation, and ANN model respectively.

Keywords: ET₀, limited data, Hargreaves, Artificial Neural Networks (ANN).
INTRODUCTION The climate has always had a very important role on human’s health. Actually, it has determined the location of agricultural production sites, recreational areas, urban development, and industrial areas. The availability of different climatic parameters permits to characterize the climate of a region, being necessary several climate data to carry out a complete study. Several studies have been based on the use of different climate data or indicators calculated by using basic climatic parameters. However, these studies usually present limitations on their development due to problems with data availability and quality (Elías and Ruiz, 1981; de León, 1988; Allen et al., 1998; Fount, 2000).

Different authors have proposed algorithms and techniques to perform data generation in different areas of knowledge (Allen et al., 1998; Allison, 2001). Some of them have tried to estimate climatic parameters, such as global radiation and reference evapotranspiration (ET0) from basic climatic data or ANN structures (Popova et al., 2006; Landeras et al., 2008; Gavilán et al., 2008).

The analysis of evapotranspiration (ET) measured at ground level has been mainly considered for the agricultural activity, in order to determine the water requirements and the supply of water in an appropriate manner (Calera, 2005). Irrigation activity uses 70% of freshwater in the world, which denotes the importance of this use (FAO, 2002). Accurate estimates of reference evapotranspiration (ET0) are required for responsible water use in irrigation (Trajkovic, 2007). ET0 determination is a key factor for water balance and crop production, water resources management, irrigation scheduling, and environmental assessment. ET0 can be obtained with indirect measurements, with high-cost micrometeorological techniques, or can also be estimated with mathematical models (Allen et al., 1991). The Food and Agriculture Organization of the United Nations (FAO) has proposed the use of the Penman-Monteith equation as the standard for estimating ET0 and for evaluating other equations (Trajkovic, 2007). However, some of the weather variables that are required by this model are often not available in irrigation. In such cases, a simple empirical Hargreaves equation is mainly used (Popova et al., 2006; Trajkovic, 2007).

The main objective of this paper is the generation of a model that estimates monthly ET0 values in Castilla-La Mancha region (Spain) from available termo-pluviometric weather stations with monthly data that can be applied to historical time series.

MATERIAL AND METHODS The proposed methodology can be summarized in the Figure 1. The first step consists in selecting the termo-pluviometric data that best estimate the ET0. From 44 complete weather stations, and with time series from 2000 to 2008, data quality tests are applied. Then, 40 weather stations and time series from 2000-2007 are selected for ANN training. With the remainder 4 stations and the year 2008 data, spatial and temporal validation is performed. Thus, the proper ANN structure can be obtained by using surface polts.

The case study The case study is located in C-LM Region (Spain) (Fig. 2), which is a semiarid region with an extension of 79,462 km². In C-LM Region, 132 weather stations are available with historical time series of monthly average, maximum, and minimum temperature data (Tav, Tmax, and Tmin), together with precipitation data, which would be useful to characterize the local climate. In addition, in this Region there are 44
complete agroclimatic weather stations (Fig. 2) which are located in irrigated areas and are included in the Agroclimatic Information Service for Irrigation (SIAR, “Servicio de Información Agroclimática para el Regadío”) network. Generation of the ANN models to estimate $ET_0$ from monthly temperature data is performed by using these data. In addition the Figure 2 shows distribution of annual average temperature in the Region.

**Calibration and validation of the Artificial Neural Network models** In the proposed methodology, Artificial Neural Networks (ANNs) are used (Schalkoff, 1997; Nabney, 2002) to generate $ET_0$ in the C-LM Region by using temperature, precipitation, and extraterrestrial solar radiation data ($Ra$) (values for $Ra$ on the 15th day of the month, Allen et al., 1998). The general structure of the used ANN is shown in Figure 3. Three layers are defined by three groups of neurons. Those layers are: input, hidden, and output nodes. Each neuron implements a local function. The input nodes implements a linear function (Eq. (1)), the hidden nodes a nonlinear function (Eq. (2)) and the output nodes implements a linear function (Eq. (4)) (Fig. 3). The equation (1) shows the activation function (weighted sum) and equation (2) the squashing function (hyperbolic tangent). In this case, in which a regression problem is tried to solve, the hidden to output layer is used the linear activation function (Eq. (3)) (Nabney, 2002).

\[
S_j = w_{0j} + \sum_{i=1}^{N} w_{ij} \cdot x_i \quad (1)
\]

\[
y_j = \tanh(S_j) \quad (2)
\]

Figure 1. Diagram of the proposed methodology.
\[ S_k = v_{0k} + \sum_{j=1}^{H} v_{jk} \cdot y_j \] (3)

Thus, the result of the ANN is described by Eq. (4). In this case (regression problem), \( Z_k = S_k \).

\[ Z_k = f(S_k) \quad \text{for } k = 1 \ldots M \] (4)

Figure 2. Location of the available and utilized complete weather stations in the Castilla-La Mancha (C-LM) Region during the time series 2000-2008. Distribution of annual average temperature.

The ANNs were trained (calibration process) under supervision using scaled conjugate gradient algorithm with the main objective of minimizing the error of the estimation.
The input layer is composed by basic climatic data such as temperature (average, maximum, and/or minimum temperature), and extraterrestrial solar radiation data. Two thousand ANN structures with different number of nodes (1 to 20) and with different number of iterations (1 to 1,000 in a range of 10) were trained with monthly data of 40 of the 44 SIAR weather stations for the period covering the years 2000 to 2007 (Figs. 2 and 4). The validation process consists on applying these 2,000 ANN structures calibrated previously to a dataset and study the generated errors. In order to decide the improved ANN structure, response surface plots are used to detect, graphically, the areas with the lowest error in both graphs. Thus the validation of the ANN models are performed from two points of view (Figs. 1 and 2): 1) spatial validation (applying the calibrated ANN structures to 4 weather stations for period 2000-2007) and 2) temporal validation (applying the calibrated ANN structures to 44 weather stations for period 2008). Data used in validation process were not used in calibration process.

Figure 3. Utilized Artificial Neural Network structure and related equations.

**Goodness of fit of the ANN models.** To determine the goodness of fit of the ANN models and Hargreaves equation (eq. (5)) different statistic parameters (Willmott, 1982) are used such as: coefficient of determination ($R^2$), Root Mean Square Error (RMSE), Relative Error (RE), and the Similarity Rate (SR) (eqs. (6) to (8). In both cases, reference values (observed variable) are the averages monthly $ET_0$ values that have been calculated by applying Penman-Monteith equation with daily date.

The Hargreaves method calculate the $ET_0$ values by using temperatures and Ra values, which depends of the latitud (eq. (5)):

$$ET_0 = 0,0023R_a(T_{av} + 17.8)\sqrt{(T_{max} - T_{min})}$$  \hspace{1cm} (5)
where: $ET_0$, is the referente evapotranspiration (mm day$^{-1}$); $R_a$, is the extraterrestrial solar radiation data (MJ m$^{-2}$ day$^{-1}$); $T_{av}$, is the monthly average temperature average ($^\circ$C); $T_{max}$, is the monthly maximum temperature average ($^\circ$C); $T_{min}$, is the monthly minimum temperature average ($^\circ$C).

$$RMSE = \left[ n^{-1} \sum (P_i - O_i)^2 \right]^{1/2}$$ \hspace{1cm} [6]

where: RMSE is the root mean square error; $n$, is the number of observations; $P_i$ and $O_i$ are the predicted and observed values, respectively.

$$RE = \left( \frac{RMSE}{O_{med}} \right) \cdot 100$$ \hspace{1cm} [7]

where: $RE$, is the relative error, estimated as a percentage of the average value of the variable; $O_{med}$, is the average value of the variable observed.

$$SR = 1 - \left[ \frac{\sum (P_i - O_i)^2}{\sum ((P_i - O_{med}) + (O_i - O_{med}))^2} \right]$$ \hspace{1cm} [8]

where: SR, is the similarity rate expressed as a relative measure of the difference between variables. (If SR = 1, there is a perfect agreement between P and O).

RESULTS

The set of variables (termo-pluviometric variables) that permits to obtain a good fitting with the $ET_0$ data by using ANN structures are $T_{max}$, $T_{min}$, and $R_a$. The response surface plots show the RMSE of all ANN structures (2,000 in this case) in function of the number of iterations and hidden nodes. As an example, Figure 4 show the response surface plots of the spatial (Fig. 4a) and temporal (Fig 4b) validation of the estimation of $ET_0$ by using as input variables $T_{max}$, $T_{min}$, and $R_a$. It can be shown that the common minimum RMSE in the estimation of $ET_0$ is obtained with a structure of between 8-16 hidden nodes and between 200-600 iterations.

![Figure 4](image_url)

Figure 4. Response surface plots of the evolution of the root mean square error (RMSE) based on Artificial Neural Network structure used for estimating the evapotranspiration of reference in Castilla-La Mancha (C-LM): a) spatial validation for 4 weather stations (time series 2000-2007), b) temporal validation for 2008 year (44 weather stations).
The selected structure for estimating ET₀ is defined by 12 nodes and 500 iterations. Table 1 summarizes the statistic parameters analyzed after the estimation of ET₀ data, for spatial and temporal validation, from average monthly Tmax, Tmin, and Ra data by using this structure. In addition, this Table shows the statistical parameters by using Hargreaves equation. The error (RMSE) is lower by using ANN vs. Hargreaves model. For spatial validation data, the RMSE values are 0.40 mm day⁻¹ and 0.46 mm day⁻¹, respectively. For temporal validation, these values are 0.42 mm day⁻¹ and 0.55 mm day⁻¹. ANN shows ERs values of 12.02% and 13.21% while in Hargreaves equation they are 13.82% and 16.99% for spatial and temporal validation data.

Table 1. Estimated monthly evapotranspiration of reference (ET₀) in Castilla-La Mancha (C-LM) by using selected ANN structure (12 nodes and 500 iterations) from averages of maximum and minimum temperatures and extraterrestrial solar radiation, and, by using Hargreaves equation. Error of spatial and temporal validation.

<table>
<thead>
<tr>
<th>Statistical parameters</th>
<th>Artificial Neural Network (ANN)</th>
<th>Hargreaves equation</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Spatial validation</td>
<td>Temporal validation</td>
</tr>
<tr>
<td>Data number</td>
<td>340</td>
<td>516</td>
</tr>
<tr>
<td>Average x</td>
<td>3.34</td>
<td>3.21</td>
</tr>
<tr>
<td>Average y</td>
<td>3.28</td>
<td>3.22</td>
</tr>
<tr>
<td>Average x /average y</td>
<td>0.98</td>
<td>1.00</td>
</tr>
<tr>
<td>R²</td>
<td>0.96</td>
<td>0.95</td>
</tr>
<tr>
<td>R² adjust</td>
<td>0.96</td>
<td>0.95</td>
</tr>
<tr>
<td>RMSE</td>
<td>0.40</td>
<td>0.42</td>
</tr>
<tr>
<td>RE</td>
<td>12.02</td>
<td>13.21</td>
</tr>
<tr>
<td>SR</td>
<td>0.99</td>
<td>0.99</td>
</tr>
</tbody>
</table>

where: x, are observed ET₀ values (mm day⁻¹); y, are predicted ET₀ values (mm day⁻¹); R², is determination coefficient; RMSE, is root mean square error (mm day⁻¹); RE, is relative error (%) and SR, is similarity rate.

**CONCLUSION.** The use of ANNs could be interesting for the generation of climate parameters from other available climatic data. For determining the optimal ANN structure in the estimation of ET₀ it is necessary to calibrate the model with part of the available data and validate the results with climatic data from stations not used in the calibration process of the ANN, considering spatial and temporal characteristics.

In C-LM Region, the estimated ET₀ data from basic temperature and Ra monthly data, were estimated with relative errors of: 1) spatial error of 12.02% 2) temporal errors of 13.21%. Hargreaves equation shows higher errors than the spatial and temporal errors obtained with the validation process of the ANN models, which are estimated in 13.82% and 16.99% respectively.

In C-LM Region, with limited data, ET₀ could be estimated by using ANN structures. This methodology shows better goodness of fit than Hargreaves equation.
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